



Probable Maximum Flood estimation using upper bounded statistical models and its effect on high return period quantiles

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3rd International Week on Risk Analysis, Dam Safety, Dam Security, and Critical Infrastructure Management Valencia, Spain, October 18-19, 2011

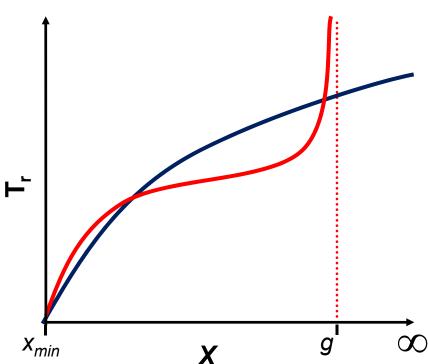




The PMF is the biggest flood physically possible at a specific catchment (Smith and Ward, 1998)

It has a physical meaning and provides an upper limit for the decision maker

It will change the cdf behaviour at medium and high return periods





- High return period quantile estimation main drawback: lack of available information about large events in a relatively short data series => increase amount of information
 - One possibility is to included palaeoflood and/or historic information. From now, Non-Systematic information:
 - Information different to the systematic record at the flow gauge station
 - No differences from the statistical point of view!



 High return period quantile estimation main drawback: lack of available information about large events in a relatively short data series => increase amount of information

 Objective: merging deterministic knowledge and statistical analysis to better estimate high return period quantiles in a framework of enough information





- Applied by Takara and Loebis (1996)
- Basic distribution function Y ~ LN2
- Slade-type transformation:

$$y = \ln\!\left(\frac{x-a}{g-x}\right)$$

Resulting pdf:
$$f(x) = \frac{g-a}{(x-a)(g-x)\sigma_y\sqrt{2\pi}} \exp\left[-\frac{1}{2}\left\{\frac{y-\mu_y}{\sigma_y}\right\}^2\right]$$

- Parameters:
 - g = upper bound (PMF)
 - a = lower bound, set to 0 to reduce the # of parameters



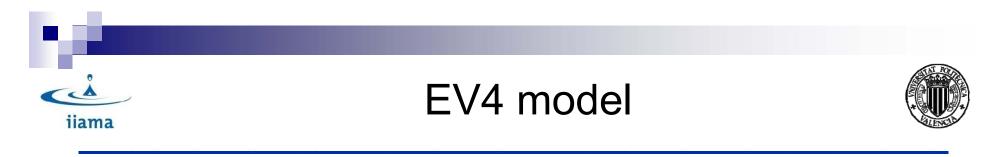


- Developed by Elíasson (1997)
- Basic distribution function: Y ~ EV1 (or Gumbel)
- Transformation: $Y = X \frac{ak^{*2}}{(a X)}$
- Resulting cdf:

$$(g-X)$$

$$F(x) = \exp\left[-\exp\left(\frac{-x}{a} + \frac{ak^*}{(g-x)} - b\right)\right]$$

- Parameters:
 - g = upper bound (PMF)
 - k* = transformation parameter = -0.5 for better results
 - a = scale and transformation parameter
 - b = location parameter



It was derived from a GEV (Takara and Tosa, 1999)

• The resulting cdf is:
$$F(x) = \exp\left[-\left\{\frac{g-x}{\nu(x-a)}\right\}^{k'}\right]$$

Parameters:

- a = lower bound (set to 0 to reduce # of parameters)
- g = upper bound (PMF)
- v = scale parameter
- k' = shape parameter





- Relative large basin: 22,000 km²
- Mediterranean torrential regime
 - Mean flow = $36 \text{ m}^3/\text{s}$
 - Mean flood = 713 m³/s
 - Q₂₀= 2,000 m³/s
 - Coefficient of variation = 2.74
 - Skewness coefficient = 5.26



 Very strong Convective Mesoscale Systems in Fall (Rigo and Llasat, 2007)

Mixed flood population (Rossi et al., 1984)

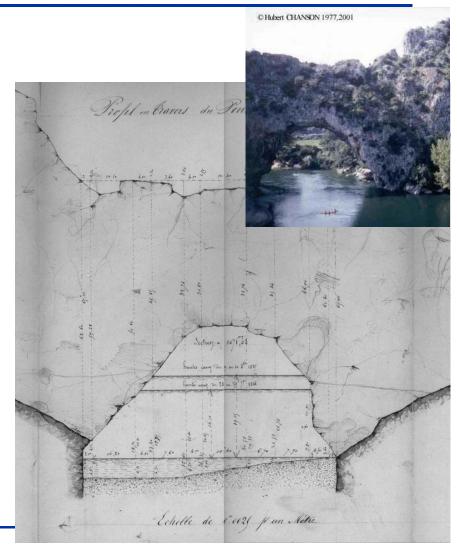




Historical information

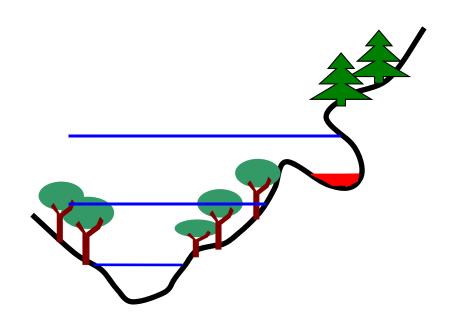
- Archives: <u>municipality</u> <u>records</u>, notary notes, engineering damage reports
- Newspapers, books, chronicles
- Maps, photographs, plans
- Building marks
- > Oral communications

Plan of the two biggest floods in XIX century floods at Pont d'Arc (Ardèche River in France)



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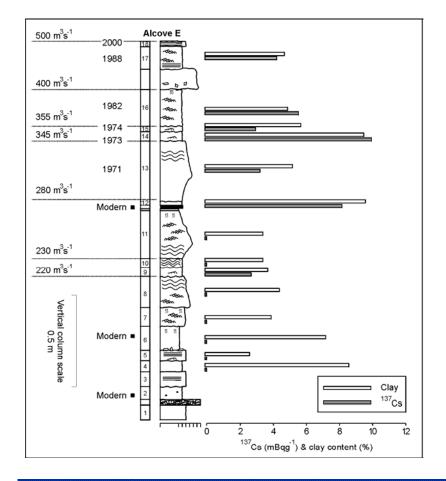


Palaeoflood information

- > Botanical evidences
- Palaeolevel indicators
 - Slackwater deposits
 - Silt marks







Palaeoflood information

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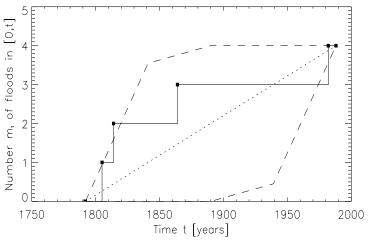
- Jucar case study with historical information (building marks): flooding of an ancient convent located in the floodplain
 - 1805
 8,400

 1814
 6,400

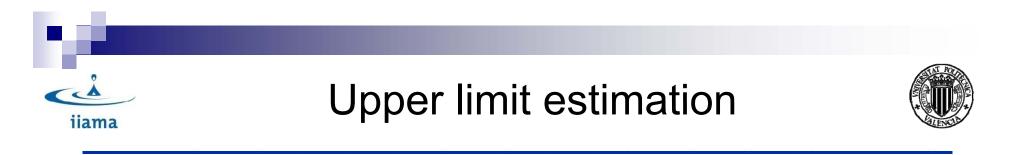
 1864
 13,000

Year

- Threshold of inundation X_H = 6,200 m³/s
- Historical period: 1792 to 1945
- Stationarity was tested and proved using the Lang Test (Lang et al., 1999)



Peak $Q (m^3/s)$



ML-PG: g is fixed at the value previously calculated (G) as the best approximation for the true unknown PMF, and the other parameters are estimated by ML method:

$$g = G$$
$$\max L(\underline{\Theta'}) \int d\theta d\theta$$

ML-C: the whole parameters set of the distribution function is estimated by the ML method, including g as another free parameter in the maximization process:

$\max L(\underline{\Theta})$





- In some combinations of distribution + type of information, ML-C estimates g as the maximum observation =>
- ML-GE: This method consists on the use of the Generic Equation to estimate g (Kijko, 2004) and the ML method for the rest of parameters:

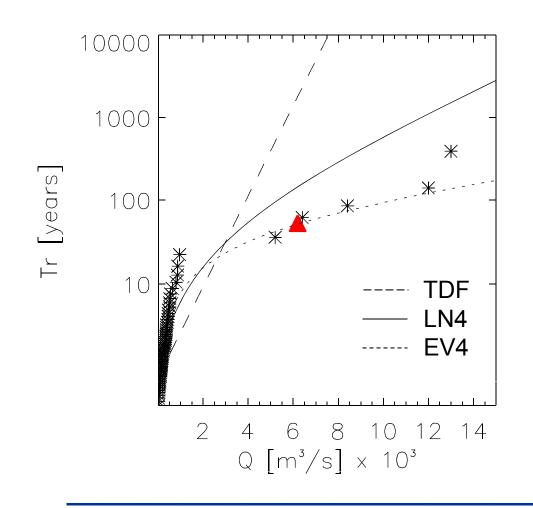
$$g = x_{\max} + \int_{-\infty}^{g} [F_X(x;\underline{\Theta'},g)]^n dx$$

$$\max L(\underline{\Theta'})$$



ML-PG (prefixed g)

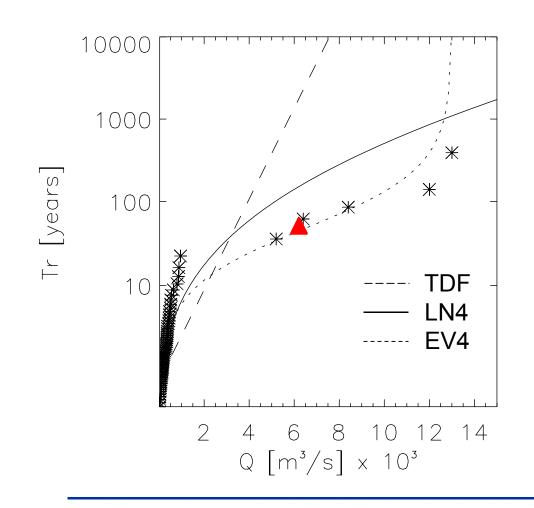




- PMF using specific discharge from upstream PMF study = 33,900 m³/s => possible overestimation
- "Dog leg" effect due to mixed populations
- Clear different approach to the upper limit
 - Slower for TDF
 - Faster for EV4



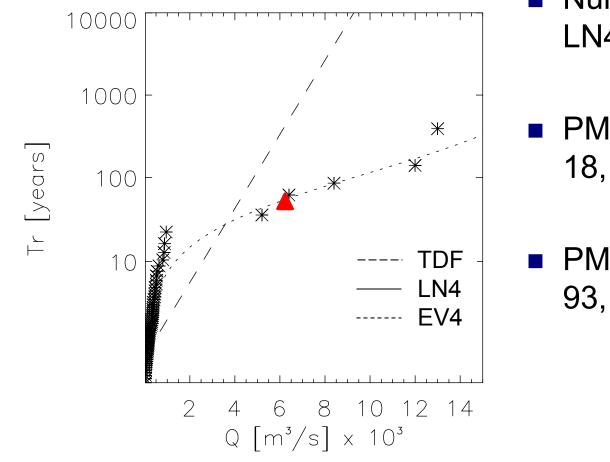




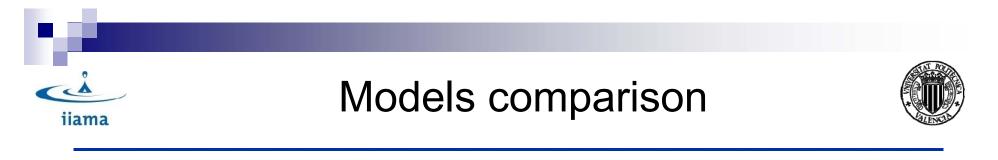
- PMF with EV4 and TDF ML-C = 13,000 m³/s (maximum observation)
- PMF with LN4 ML-C = 93,300 m³/s



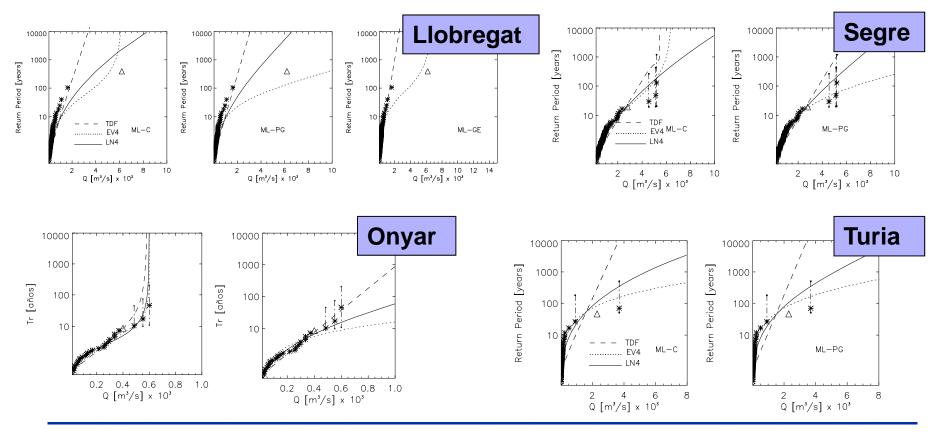




- Numerical problems with LN4
- PMF with EV4 ML-GE = 18,100 m³/s
- PMF with TDF ML-GE = 93,100 m³/s



 EV4 better performance with high skewnees coefficient, as it was pointed out by Takara and Tosa (1999)

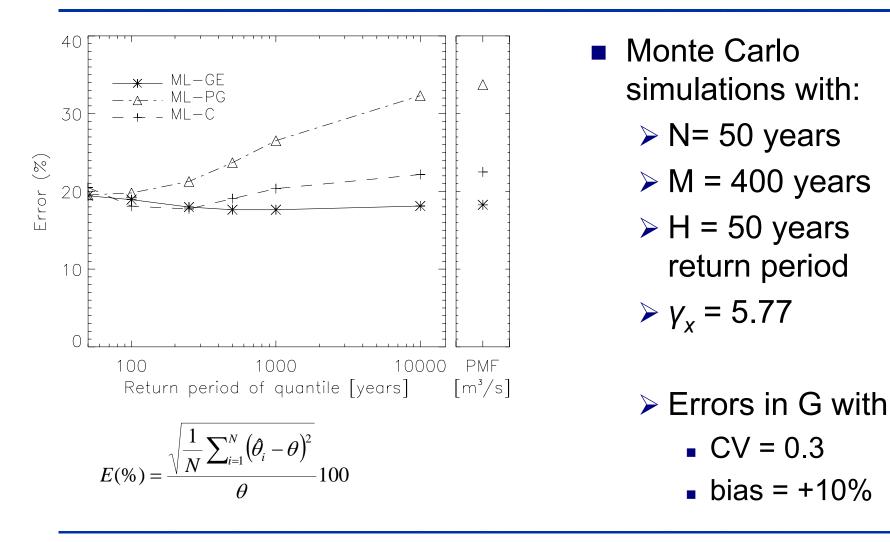


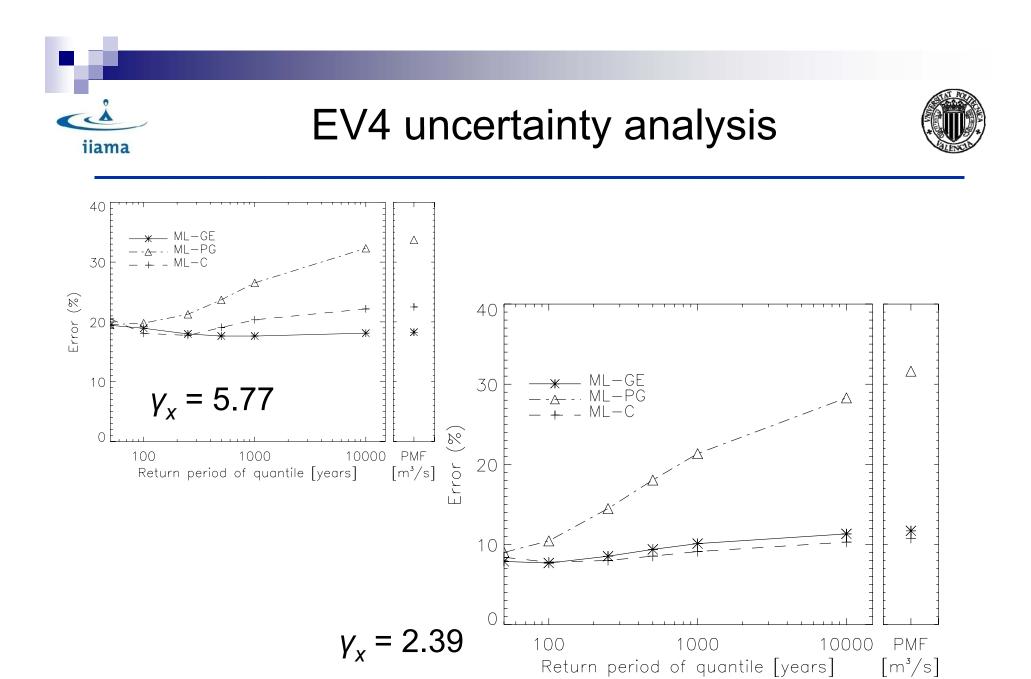
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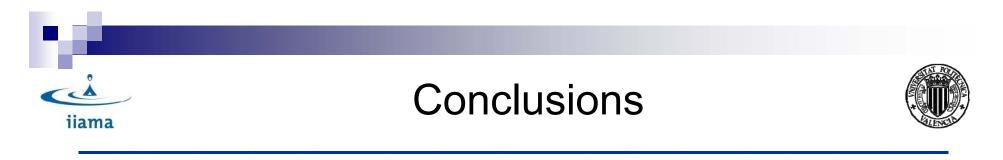


EV4 uncertainty analysis

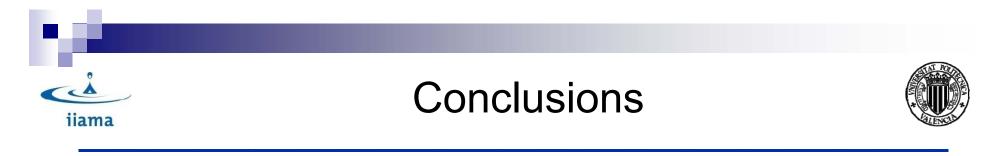








- Why not explore and use upper bounded distribution functions?
 - There is an upper limit
 - Upper bounded and unbounded distributions behave different at medium and high return periods



- Why not explore and use upper bounded distribution functions?
- The upper limit represents the PMF and either can be:
 Causal information expansion (Merz and Blöschl, 2008), if it is fixed a priori
 - Within a temporal information expansion framework (Merz and Blöschl, 2008), considered as one more parameter to be estimated in the statistical model fitting



Program **AFINS** available at http://lluvia.dihma.upv.es

Poster: High return period annual maximum reservoir water level quantiles estimation using synthetic generated flood events

Many thanks for your attention!

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