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Hydrological and sediment model calibration at ungauged basins using check dam stratigraphy as proxy data in a Mediterranean catchment

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- ❑ *Problem:* hydrological modelling in Mediterranean semi-arid zones is limited by **data availability**:
 - **Good** data availability: rainfall, temperature, land use, ...
 - **Scarce** data: water & sediment discharge, soil properties, ...

- ❑ *Aim of the work:* calibration and validation of a hydrological model in a semi-arid catchment without discharge data.
 - 1 – simplifying the model by making realistic hypothesis;
 - 2 – using proxy data: **check dam sedimentation volume**;
 - 3 – using multidisciplinary techniques:
 - **stratigraphical analysis**
 - **hydrological modelling**
 - **reservoir sedimentation modelling**

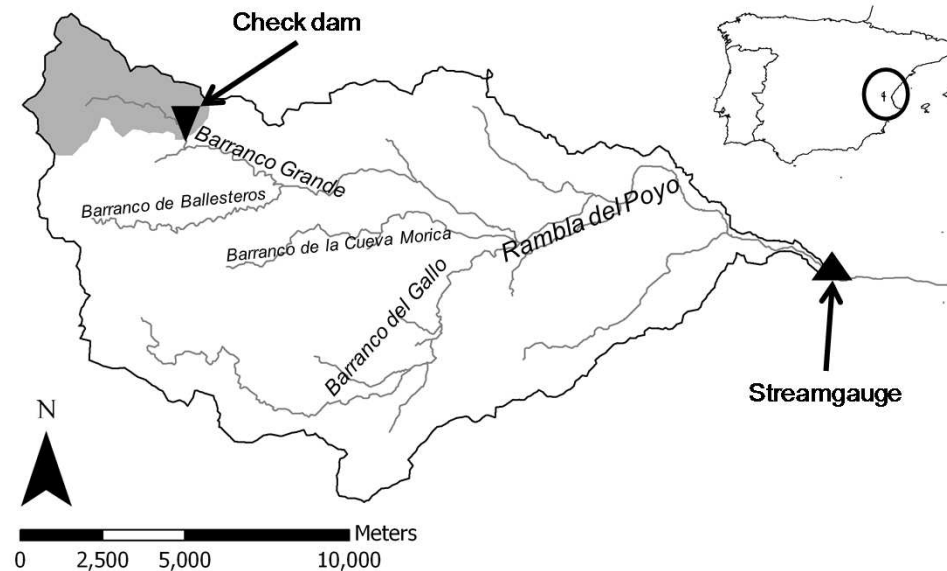


□ Methodology:

- 1 – Reconstruction of the **depositional record** of a check dam infill deposit by means of a stratigraphical description;
- 2 – **Calibration and validation** of a hydrological and sediment model using the reconstructed depositional record;
- 3 – **Verification** of the model performance by comparing simulated and observed water discharge at the streamgauge station;
- 4 – **Sensitivity analysis** (Monte Carlo simulation).



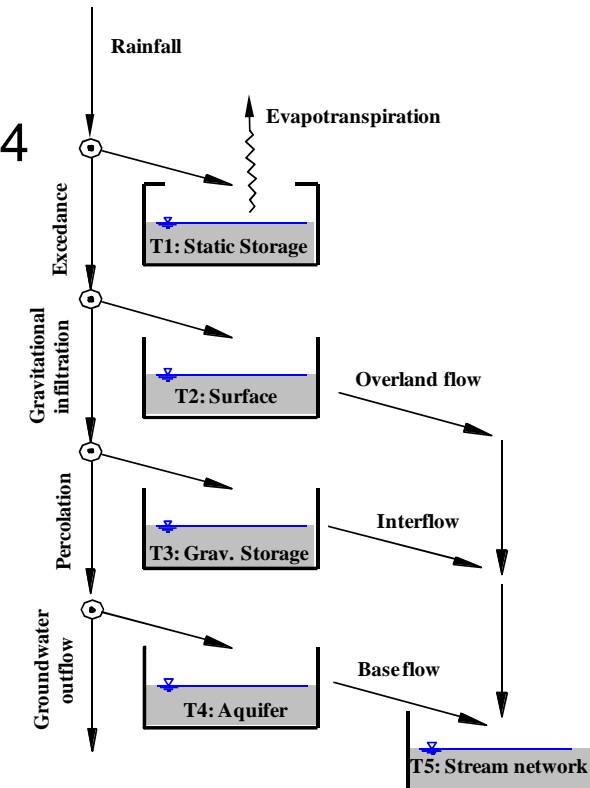
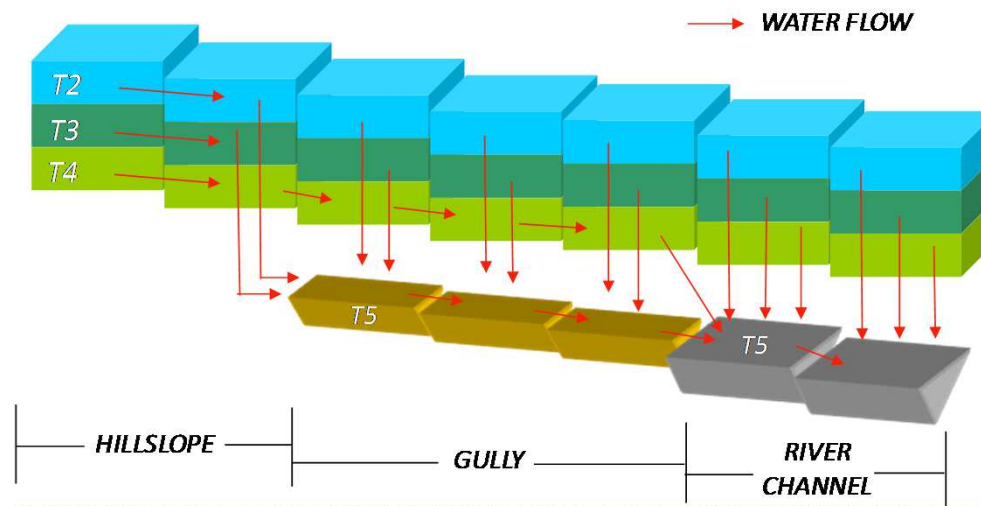
- Rambla del Poyo catchment (Valencia, SE Spain)
 - Semi-arid climate (rainfall = 450 mm/year; $ET_0 = 1,100$ mm/year)
 - Geology dominated by limestone
 - Shrubland cover (*matorral*)
 - Studied catchment: 12.9 km²
 - Streamgauge catchment: 184 km²





□ TETIS model: hydrological sub-model

- Developed in the TU of Valencia since 1994
- Distributed and conceptual (tank structure) model, with physically based parameters
- Reproduction of hydrological cycle spatial variability
- It uses all spatial information available



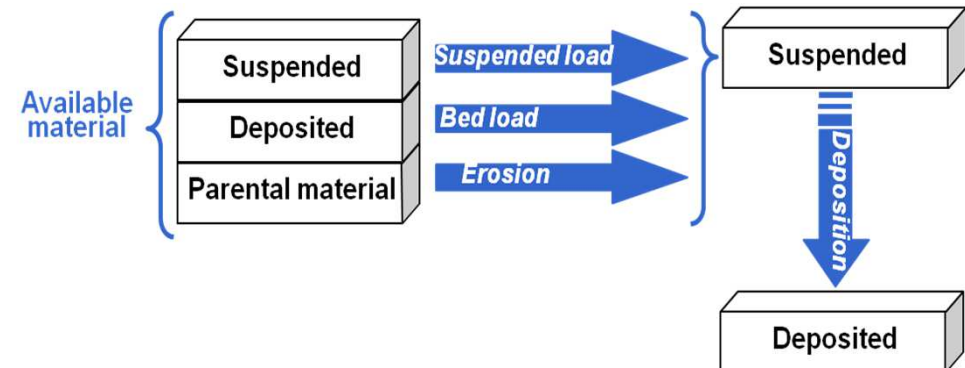


□ TETIS model: sediment sub-model

- Integration of CASC2D-SED (Julien and Rojas, 2002) in TETIS
- Balance between water transport capacity and sediment availability
- Hillslope transport capacity: modified Kilinc – Richardson (1) equation (Julien, 1995)
- Gully and channel transport: Engelund – Hansen equation (2)

$$(1) \quad Q_h = \frac{1}{\gamma_s} W \alpha S_o^{1.66} \left(\frac{Q}{W} \right)^{2.035} \frac{K}{0.15} C P$$

$$(2) \quad C_{w,i} = \beta \left(\frac{G}{G-1} \right) \frac{V S_f}{\sqrt{(G-1) g d_i}} \sqrt{\frac{R_h S_f}{(G-1) d_i}}$$



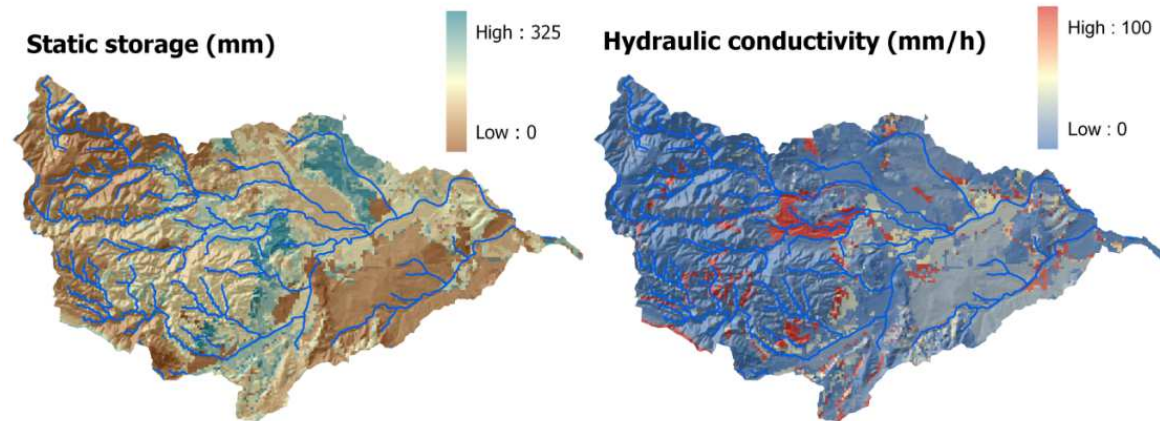
- Reservoir sedimentation: **STEP model** (Verstraeten and Poesen, 2001)



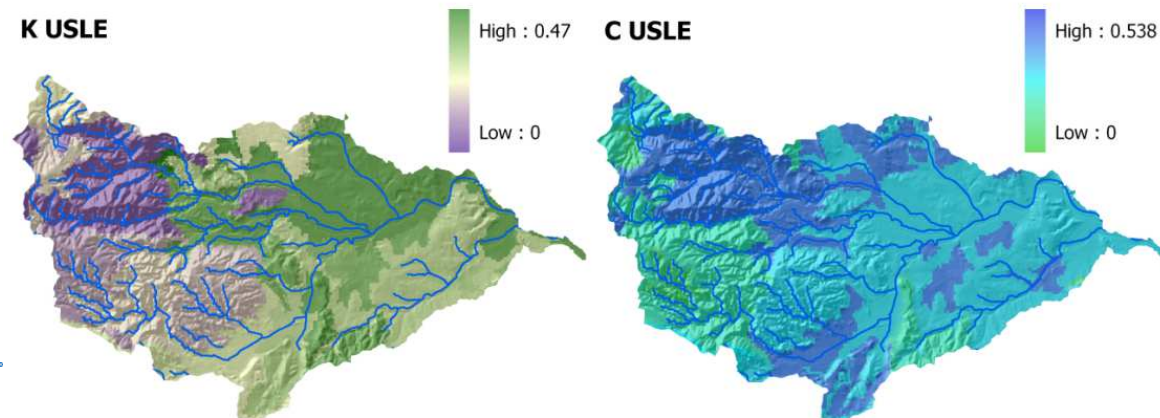
The model parameters

□ Model parameters:

- Soil hydrological properties (static storage, hydraulic conductivity, ...);



- Sediment production properties (C, K and P factor of USLE and soil texture).

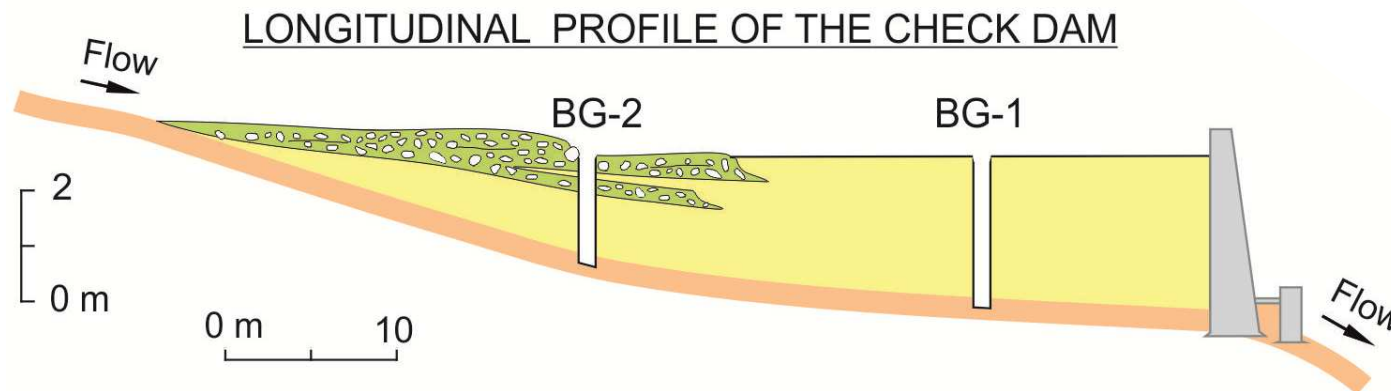


□ Check dam infill volume

- Many small check dams (2 – 10 m tall) were built in Spanish Mediterranean during 90s as sediment traps;
- A partially filled check dam was chosen:
 - Height: 4.5 m;
 - Catchment: 12.9 km²;
 - Capacity: 3000 m³;
 - Total infill ~ 1400 m³.



- ❑ GPS survey for infill volume estimation
- ❑ Two 10 x 2.5 m trenches dug across the dam infill
- ❑ Detailed stratigraphic panels with 1 m mesh
 - detection of alluvial layers deposited by different floods (the separation is defined by a break in deposition)

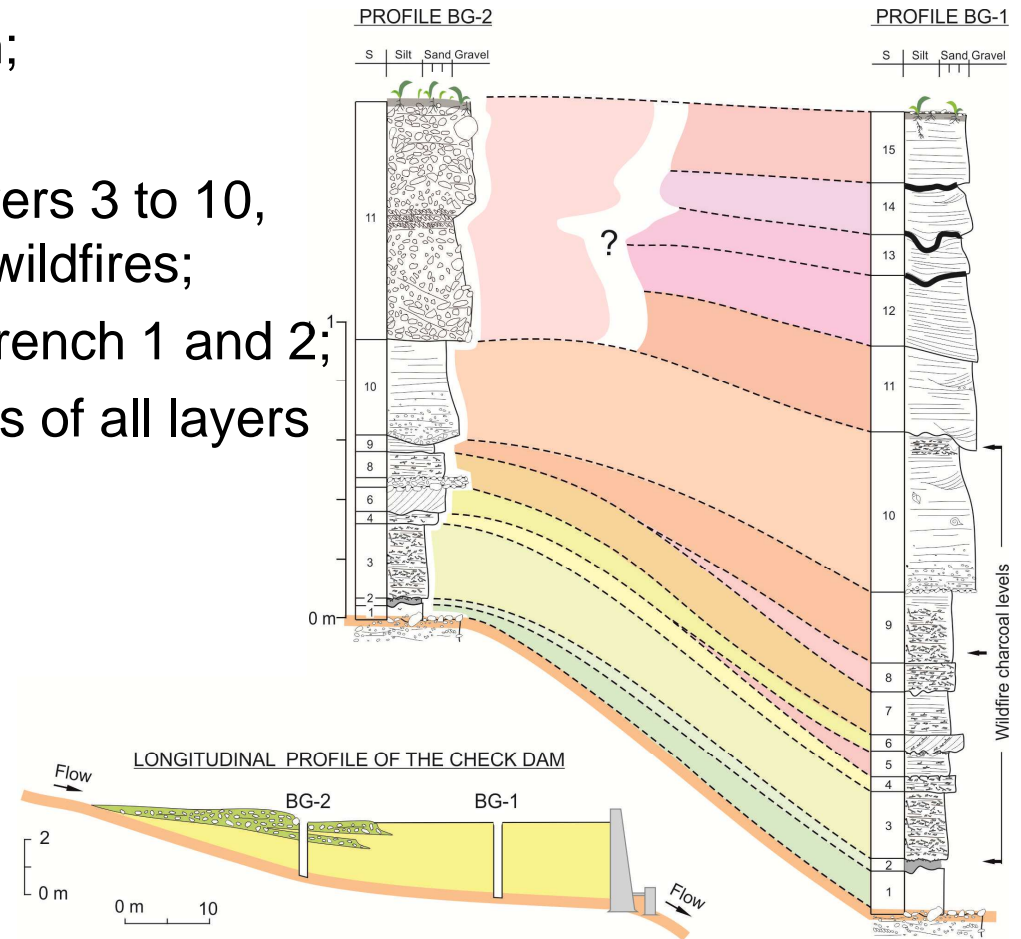




Infill volume estimation

□ Stratigraphical description

- 2.5 m sediment column;
- 15 layers (flood units);
- rests of **charcoal** in layers 3 to 10, due to 1994 and 2000 wildfires;
- 8 layers found in both trench 1 and 2;
- granulometrical analysis of all layers (sandy sediments);



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HESS Discussion



□ Two methodologies:

1 – **wedge approach**: every layer volume was calculated as if each flood unit had a wedge shape.

2 - **proportional approach**: by subtracting to the actual deposits the average layer depth.

Flood unit	Volume i) (m ³)	Volume ii) (m ³)
1	34	38
2	8	28
3	172	78
4	10	27
5	14	18
6	55	18
7	22	11
8	20	41
9	195	96
10	153	233
11	75	110
12	8	11
13	37	46
14	30	23
15	18	22
surface	582	448
tot	1434	1248



- ❑ The model need to be **simplified**

- ❑ Some hypothesis(confirmed by field observations):
 - Hortonian flow
 - Very little interflow
 - No base flow

- ❑ Parameters to calibrate (5 most influential parameters):
 - Upper soil static storage
 - Upper soil vertical hydraulic conductivity
 - Upper soil horizontal hydraulic conductivity
 - Routing correction coefficient
 - α : Kilinc – Richardson sediment production coefficient

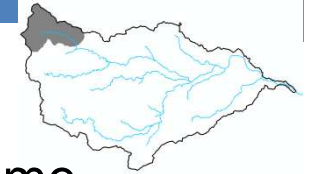


□ Dating:

- Using charcoal and knowing wildfires dates (1994 and 2000);
- Model results help dating;

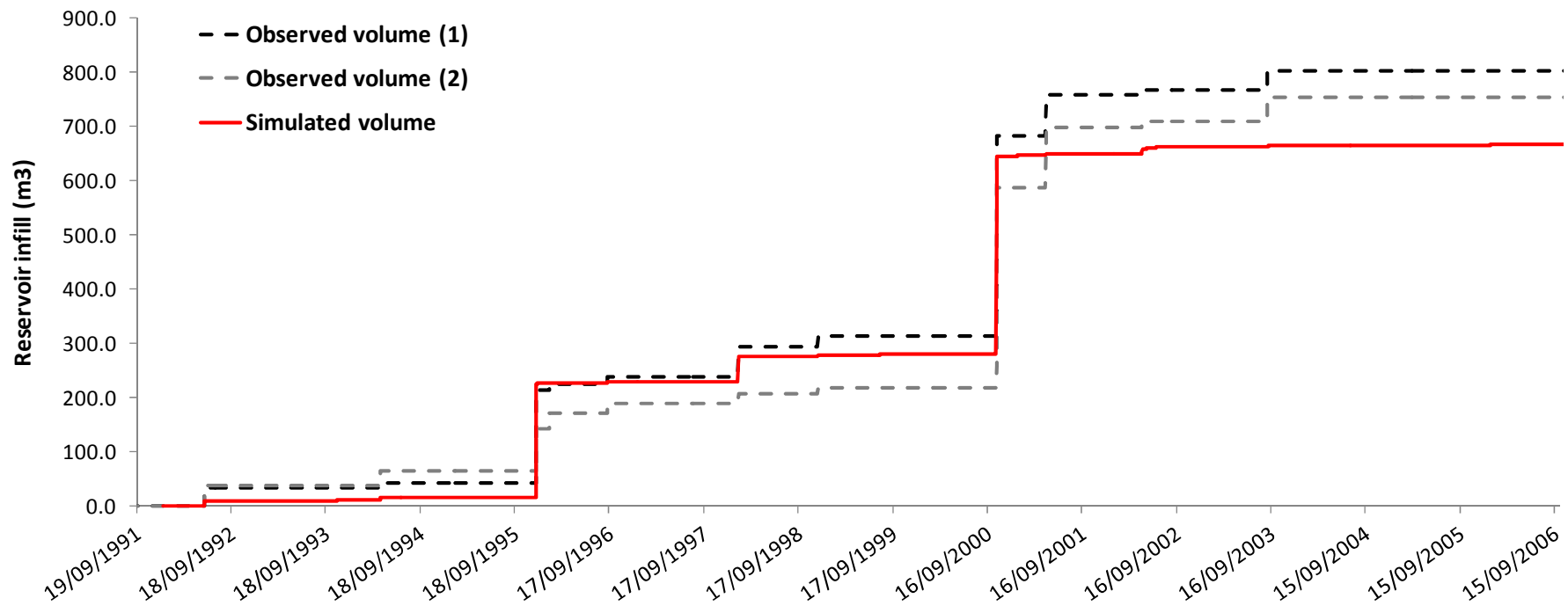
□ Calibration

- The model is calibrated simulating the reservoir observed deposited volume of the **October 2000 event** (the most extreme event in the historical series) with a **daily Δt** ;
- The October 2000 deposited layer is formed by flood units 8 + 9 +10;
- Deposited volume is $\sim 370 \text{ m}^3$;



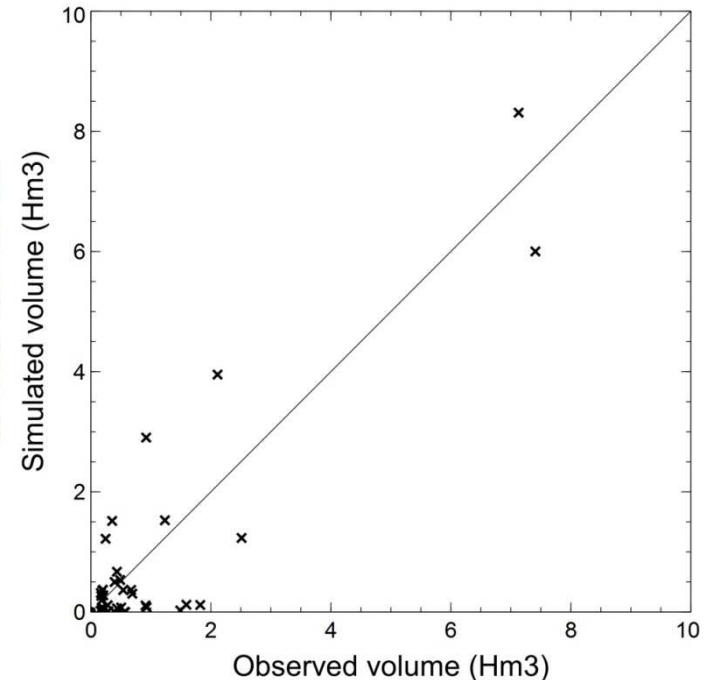
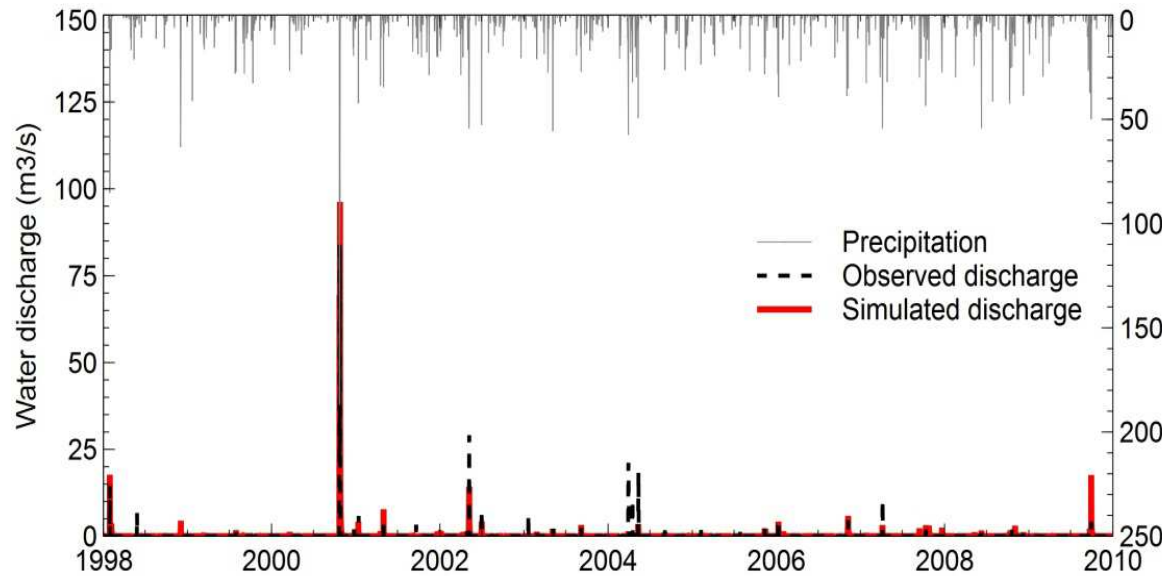
□ Sediment model validation

- The model is validated vs observed reservoir infill volume (reconstructed from stratigraphical description) – 12.9 km²
- **Feedback** process: model results help dating sediment layers



□ Water discharge validation

- The model is validated vs water discharge from the Rambla del Poyo streamgauge (184 km²)

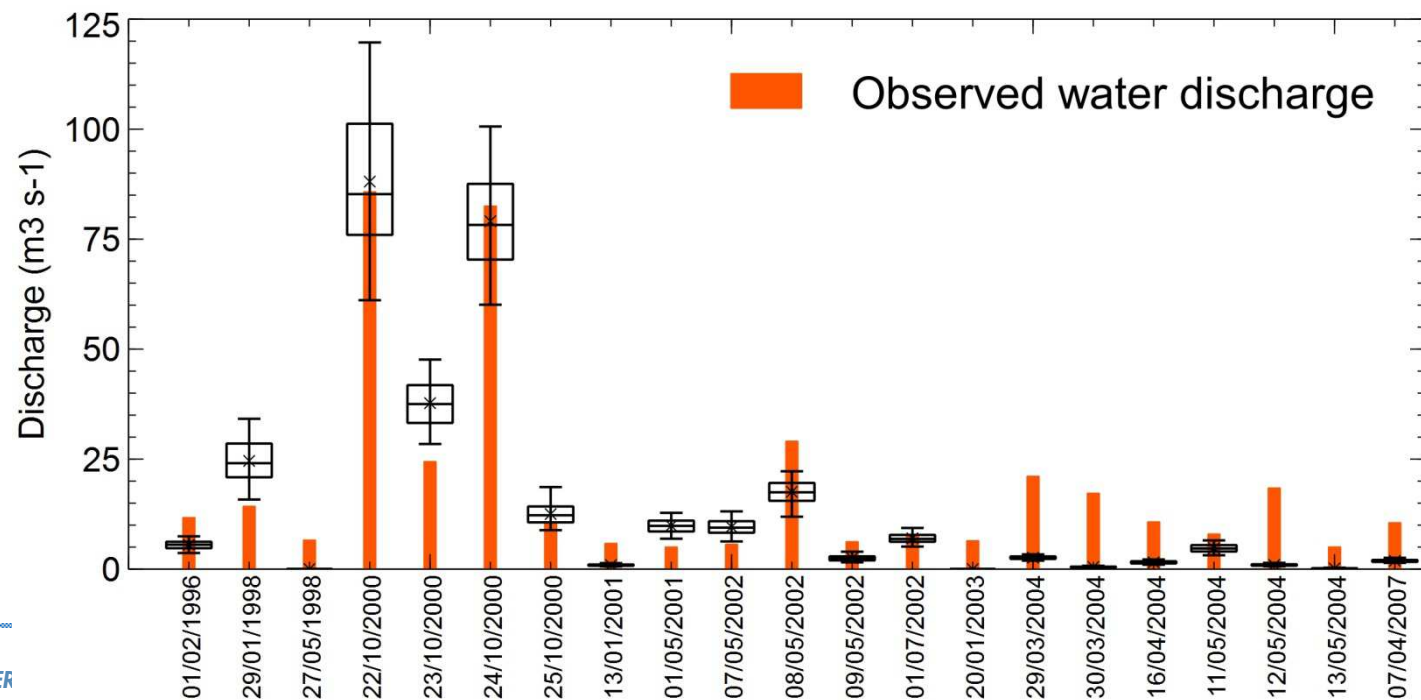


NSE index = 0.86



□ Monte Carlo simulation

- 50,000 simulations with random values of the 4 most influent parameters;
- The best 100 simulations in terms of total infill **volume error** at the check dam are tested at the stream gauge in terms of **NSE**.





- ❑ Sediment proxy data help constrain water cycle model calibration (**transfer of information** from sediment cycle to water cycle);
- ❑ **Multidisciplinarity**: coupling hydrological modelling and palaeohydrological techniques for improving catchment knowledge;
- ❑ **Small data requirement**: rainfall and temperature, soil data, land use and partially filled check dams;
- ❑ **Generalization**: this technique can be used in almost all Mediterranean small and medium size catchment



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Thanks for your attention!

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